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Early Online Release: This preliminary version has been accepted for publication in *Bulletin of the American Meteorological Society*, may be fully cited, and has been assigned DOI 10.1175/BAMS-D-19-0037.1. The final typeset copyedited article will replace the EOR at the above DOI when it is published.

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231	Abstract
232	Weather and climate variations on subseasonal to decadal timescales can have enormous
233	social, economic and environmental impacts, making skillful predictions on these timescales a
234	valuable tool for decision makers. As such, there is a growing interest in the scientific,
235	operational, and applications communities in developing forecasts to improve our
236	foreknowledge of extreme events. On subseasonal to seasonal (S2S) timescales, these include
237	high-impact meteorological events such as tropical cyclones, extratropical storms, floods,
238	droughts, and heat and cold waves. On seasonal to decadal (S2D) timescales, while the focus
239	broadly remains similar, (e.g., on precipitation, surface and upper ocean temperatures and their
240	effects on the probabilities of high-impact meteorological events), understanding the roles of
241	internal and externally-forced variability such as anthropogenic warming in forecasts also
242	becomes important.
243	
244	The S2S and S2D communities share common scientific and technical challenges. These include
245	forecast initialization and ensemble generation; initialization shock and drift; understanding the
246	onset of model systematic errors; bias correction, calibration, and forecast quality assessment;
247	model resolution; atmosphere-ocean coupling; sources and expectations for predictability; and

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248	linking research, operational forecasting, and end user needs. In September 2018 a coordinated
249	pair of international conferences, framed by the above challenges, was organized jointly by the
250	World Climate Research Programme (WCRP) and the World Weather Research Programme
251	(WWRP). These conferences surveyed the state of S2S and S2D prediction, ongoing research,
252	and future needs, providing an ideal basis for synthesizing current and emerging developments
253	in these areas that promise to enhance future operational services. This article provides such a
254	synthesis.
255	
256	Capsule
257	Climate prediction on subseasonal to decadal time scales is a rapidly advancing field that is
258	synthesizing improvements in climate process understanding and modeling to improve and
259	expand operational services worldwide.
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270 [Introductory text]

271 Beyond the tremendous progress in weather forecasting witnessed in recent decades (Bauer et al. 2015), predictive capabilities have expanded, increasingly seamlessly, to encompass climate 272 273 on subseasonal to decadal time scales (Fig. 1 and Kirtman et al. 2013). These advances have 274 been enabled by better observations, data assimilation schemes, and models originating both from the weather prediction and long term climate simulation communities, together with 275 276 increased computational power supporting progressively higher resolution and larger 277 ensembles that allow uncertainties to be better estimated and in some cases reduced. 278 International efforts under the auspices of the World Weather Research Programme (WWRP) 279 280 and World Climate Research Programme (WCRP) have helped drive this progress through coordinated research to improve the accuracy and utilization of weather and climate 281 282 predictions. Community research efforts under the WCRP led initially to climate predictions one to two seasons ahead becoming part of the World Meteorological Organization (WMO) 283 operational infrastructure (Graham et al. 2011). More recently a joint WWRP and WCRP 284 Subseasonal to Seasonal Prediction Project has started tackling the so-called weather-climate 285 prediction desert from two weeks to a season (Robertson et al. 2018; Mariotti et al. 2018), 286 287 aiming to underpin new WMO operations on those time scales (Vitart et al. 2017), and the 288 NOAA-led SubX project has similar aims (Pegion et al. 2019). At longer ranges, WCRP-enabled research has quantified predictability from a year to a decade, and corresponding WMO 289 operational infrastructure for annual-to-decadal climate prediction is now in place (World 290 291 Meteorological Organization 2018; Kushnir et al. 2019).

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Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

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293	As each of these efforts has progressed it has become increasingly apparent that common
294	challenges exist across predictive time scales. These include understanding and adequately
295	representing in models processes that give rise to predictability in the Earth system, consisting
296	of the physical climate system—atmosphere, ocean, land and sea ice—together with associated
297	biogeochemical cycling, especially of carbon (upper part of Fig. 1); capturing and
298	communicating inherent uncertainties caused by the chaotic nature of weather and climate;
299	correcting for and reducing imperfections in models that may systematically degrade forecast
300	quality; and providing forecast information in a form that is applicable to decision making. At
301	the same time, opportunities for usefully predicting elements of the Earth system beyond long-
302	term means of standard meteorological variables, including land, ocean and sea ice properties
303	and risks of weather extremes, have come into focus. The ultimate collective endeavor is to
304	improve the prediction of the spatial-temporal continuum connecting weather to climate
305	through a coordinated, seamless and integrated Earth system approach for the benefit of
306	society.
307	
308	In September 2018, international conferences ¹ on subseasonal to seasonal prediction (S2S,

309 encompassing forecast ranges from two weeks to a season) and seasonal to decadal prediction (S2D,

encompassing ranges longer than a season, up to a decade) together with cross-cutting plenary

¹ The Second International Conference on Subseasonal to Seasonal Prediction (S2S) and Second International Conference on Seasonal to Decadal Prediction (S2D) were held 17-21 September 2018 at NCAR facilities in Boulder Colorado. These coordinated meetings involved 347 participants, including 92 early career scientists, from 38 countries, with a total of 368 oral and poster presentations. Further information including a complete list of contributions can be found at <u>https://www.wcrp-climate.org/s2s-s2d-2018-home</u>.

311	sessions were convened jointly by WWRP and WCRP. This represented a confluence of research
312	and operational climate prediction expertise and knowledge exchange across prediction time
313	scales that was unprecedented in scope. Selected outcomes, organized by themes
314	encompassing the challenges outlined above, are synthesized in this article.
315	
316	Mechanisms of predictability.
317	Subseasonal to Seasonal
318	A major source of S2S predictability is the organization of tropical convection by the Madden
319	Julian Oscillation, or MJO (Woolnough, 2019), which is predicted skillfully by S2S project models
320	up to 3-4 weeks ahead (Vitart 2017). The MJO has worldwide impacts that depend on its
321	amplitude and phase, including modulation of tropical cyclone activity (Lee et al. 2018; Zhao et
322	al. 2019) and extratropical phenomena such as the East Asian summer monsoon (Li et al. 2018).
323	The associated tropical-extratropical teleconnections (Lin et al. 2019) impart S2S forecast skill
324	for many of these extratropical phenomena including Euro-Atlantic weather regimes, position
325	of the jet stream, atmospheric rivers (DeFlorio et al. 2019), and hail/tornado activity (Baggett et
326	al. 2018). However, good representations of the basic state both in the tropics and extratropics,
327	as well as tropical air-sea interactions and atmospheric convection (e.g., Yoo et al. 2015), are
328	necessary for these teleconnections to be correctly simulated by general circulation models
329	(Henderson et al. 2017).

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Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

S2S predictability also derives from the stratosphere through its relatively long time scales of 331 332 variability² and lagged influences on the troposphere (Kidston et al. 2015). Interactions 333 between the stratosphere and the troposphere from the tropics to the extratropics thus provide a promising source of S2S prediction skill (Butler et al. 2019). For example, in the winter 334 335 Northern Hemisphere stratosphere the climatological westerly polar vortex exhibits extremes 336 in variability, including sudden stratospheric warmings (SSWs) that are driven largely by Rossby waves from the troposphere. SSWs have lagged impacts on sea level pressure, surface 337 338 temperature and precipitation, including pronounced tendencies for cold anomalies over 339 northern Eurasia and warm anomalies over northeastern North America (e.g., Sigmond et al. 2013). Initializing forecasts during extreme stratospheric events provides increases in prediction 340 341 skill of surface climate in such regions up to 3-6 weeks later (Domeisen et al. 2019c). However, the predictability of specific extreme stratospheric events is limited, ranging from a few days to 342 343 about two weeks (Fig. 2) for different SSWs (Karpechko 2018; Taguchi 2018, Domeisen et al. 2019a), although models show evidence of under-confident forecasts in the stratosphere on 344 S2S timescales (O'Reilly et al. 2019). Outstanding questions remain about the mechanisms of 345 346 stratosphere-troposphere coupling processes, in particular on the causes, variability, and trends 347 for the occurrence of SSW events (Ayarzaguena et al. 2018; Simpson et al. 2018) and why not all SSW events have similar downward effects (e.g., Garfinkel et al. 2013, Maycock & Hitchcock, 348 349 2015). In addition, further research is needed to assess the degree to which prediction models capture both the stratospheric variability and coupling processes. 350

² Including the quasi-biennial oscillation (QBO) of the tropical stratosphere, whose influences span a range of time scales and are addressed in the "Time scale interactions" subsection.

352	Among atmosphere-surface influences, land-atmosphere interactions have their greatest
353	impact on subseasonal time scales in forecasts where land is initialized (Dirmeyer et al. 2018a),
354	but also can contribute skill on weather prediction and multi-month time scales (Dirmeyer and
355	Halder 2016, 2017). The most broadly impactful land attribute is soil moisture (Koster et al.
356	2004, 2016), but anomalies in soil temperature (Y. Zhang et al. 2019; Yang et al. 2019), snow
357	cover (Jeong et al. 2012; Orsolini et al. 2013), and vegetation states (Williams et al. 2016) can all
358	have significant impacts. A number of recent studies have focused on non-local impacts of land
359	surface anomalies, showing for example that soil moisture anomalies can exert remote as well
360	as local influences in boreal summer through driving of quasi-stationary Rossby waves and
361	associated circulation anomalies (e.g., Teng et al. 2019; Wang et al. 2019). In addition, land
362	surface and subsurface temperatures in spring may exert delayed downstream influences on
363	precipitation (Xue et al. 2018), and evapotranspiration may remotely influence precipitation
364	over land (Wei and Dirmeyer 2019).
365	
366	Atmosphere-ocean interactions, fundamental for S2D predictability, can also be influential on
367	S2S time scales. For example submonthly prediction skills for precipitation and temperature are
368	enhanced over certain land areas including parts of Australia, the Maritime Continent and the
369	contiguous United States when tropical sea surface temperature (SST) anomalies associated
370	with El Niño Southern Oscillation (ENSO) are present (Hudson et al. 2011; Li and Robertson
371	2015; DelSole et al. 2017). Extratropical SST anomalies also can impart S2S skill through
372	teleconnections, as shown for example by McKinnon et al. (2016) who identified a SST anomaly

pattern in the mid-latitude North Pacific that tends to precede heat waves and rainfall deficits
in the eastern United States by up to 50 days.

375

Sea ice strongly influences surface fluxes and lower atmospheric temperatures particularly in the marginal ice zone, and provides a source of S2S predictability for polar and possibly midlatitude regions (Chevallier et al. 2019). This motivates the development of S2S forecasts for sea ice, which thus far have shown significant, albeit region-dependent skill for predicting intraseasonal Arctic sea ice variability (Liu et al. 2018, Zampieri et al. 2018).

381

382 Seasonal to decadal

383 A primary general source of S2D atmospheric predictability is remote influences from a variety of teleconnections (e.g., Yuan et al. 2018; Ruprich-Robert et al. 2018; Beverley et al. 2019. 384 385 Teleconnections associated with anomalous atmospheric circulation patterns arise from changes to the Walker circulation usually driven by anomalous zonal SST gradients (Cai et al. 386 387 2019), and changes to the Hadley circulation usually driven by anomalous meridional SST gradients, especially interhemispheric differences (Kang et al. 2018). These influences impact 388 tropical cyclones and rainfall, whereas anomalous upper level divergence due to tropical rainfall 389 390 anomalies leads to Rossby waves that impact the extratropics (Scaife et al. 2017; O'Reilly et al. 391 2018). Besides giving rise to atmosphere-ocean interactions that alter the atmospheric circulation, SST anomalies can induce low-level temperature and moisture anomalies that are 392 393 advected elsewhere by climatological winds (Dunstone et al. 2018).

S2D atmospheric predictability arising from teleconnections requires that SST anomalies be 395 396 predictable. On seasonal timescales, tropical SST anomalies are dominated by ENSO (Yang et al. 2018), though there is some independent variability in the tropical Atlantic and Indian Oceans 397 that also drives teleconnections (e.g., Nnamchi et al. 2015; Lim et al. 2016). The impacts of 398 399 ENSO are sensitive to ENSO diversity (Capotondi et al. 2015), including the longitude at which maximum SST anomalies occur (Yeh et al. 2018; Patricola et al. 2018). ENSO SST anomalies are 400 largely predictable out to a year particularly in winter and early spring (Barnston et al. 2017), 401 402 whereas predictability may extend to two years for some La Niña events (Di Nezio et al. 2017), 403 and to 1 ½ to two years for certain El Niño events (Luo et al. 2008).

404

405 Decadal SST variability occurs in both the Atlantic and Pacific oceans, often referred to as Atlantic Multidecadal Variability (AMV) and Pacific Decadal Variability (PDV), e.g. Kushnir et al. 406 407 (2019). The causes of AMV are not fully understood, especially the relative roles of internal variability and external forcing from aerosols. However, AMV is modulated to some extent by 408 409 the oceanic Atlantic Meridional Overturning Circulation (Yeager and Robson 2017), which together with the North Atlantic subpolar gyre is influenced by deep ocean density anomalies 410 particularly in the Labrador Sea (Robson et al. 2016); these influences contribute to the 411 412 especially high multi-year predictability in the North Atlantic (Buckley et al. 2019). AMV couples 413 to the Hadley circulation, affecting hurricanes and Sahel rainfall as illustrated in Fig. 3 (Sheen et al. 2017), and can initiate atmospheric Rossby waves with remote influences including 414 temperatures in parts of China (Monerie et al. 2018). AMV can influence PDV (Ruprich-Robert 415 416 et al. 2017), and vice-versa. PDV may also be influenced by off-equatorial heat content

417	anomalies in the western Pacific Ocean (Meehl et al. 2016). Decadal variability of deep
418	convection in the Southern Ocean influences temperatures in that region, potentially explaining
419	recent increases in Antarctic sea ice (L. Zhang et al. 2019).
420	
421	S2D atmospheric predictability also arises from longer time scale processes over land, mainly
422	involving soil moisture (Chikamoto et al. 2017; Ardilouze et al. 2019) and vegetation (Weiss et
423	al. 2014; Bellucci et al. 2015). These highlight the need for land surface initialization
424	(Prodhomme et al. 2016a) and realistic vegetation models (Alessandri et al. 2017).
425	
426	An additional source of S2D predictability is variations in radiative forcing, which provide
427	significant skill on multi-year timescales (Smith et al. 2019). Much of this skill arises from
428	changes in greenhouse gases, but anthropogenic aerosols may force decadal variations in AMV
429	(Booth et al. 2012) and PDV (Smith et al. 2016; Takahashi and Watanabe 2016). Solar variability
430	(Misios et al. 2019), and volcanic eruptions (Menegoz et al. 2018) including through their
431	influence on ENSO (Khodri et al. 2017; Wang et al. 2018) and possibly AMV and the North
432	Atlantic Oscillation (NAO; Swingedouw et al. 2017) affect climate on seasonal to decadal
433	timescales and are potentially important sources of predictability. However, the relative roles
434	of external radiative forcing and internal variability (W. Kim et al. 2018) continue to be
435	explored.
436	

437 Time scale interactions

The Quasi-biennial Oscillation (QBO) is a downward-propagating ~28-month oscillation of 438 439 easterly and westerly zonal jets in the tropical stratosphere, driven by upward equatorial waves from the troposphere (e.g., Kim and Chun 2015). In addition to having high predictability and 440 441 some teleconnected influence on winter surface climate (e.g., Scaife et al. 2014a), the QBO 442 modulates the amplitude, persistence, and rate of propagation of the boreal wintertime MJO 443 (Fig. 4) through its impact on tropical convection via changes in static stability near the tropopause (Yoo and Son 2016, Nishimoto and Yoden 2017). MJO amplitude is better predicted 444 445 at longer leads during the easterly phase of the QBO (Marshall et al. 2017), likely as a result of longer persistence of the MJO rather than its greater initial amplitude (Lim et al. 2019). 446 447 448 The modulation of SSW probability of occurrence by tropical sources of variability, such as the QBO, ENSO, or MJO, may extend probabilistic predictability of stratospheric variability to a few 449 450 months or longer if these relationships can be adequately captured by prediction models 451 (Garfinkel & Schwartz 2017; Garfinkel et al. 2018; Domeisen et al. 2019a,b). 452 There is increasing evidence of additional interactions between various sources of S2S and S2D 453 454 predictability across time scales. One example is that seasonal time scale variations in ENSO 455 modulate the MJO (Chen et al. 2016) and its impact on the NAO (Lee et al. 2019) with 456 consequent influences on weather over remote regions. Another is that ENSO teleconnection to the extratropics has varied over multi-decadal time scales spanning the past 100+ years 457 458 (O'Reilly 2018), possibly modulating ability to predict the NAO (Weishiemer et al. 2019),

although sampling variability can also give rise to such long-term changes in teleconnections(Yun and Timmermann 2018).

461

462 Modelling issues.

463 Subseasonal to Seasonal

Because S2S operational prediction is a relatively new enterprise, considerable efforts focusing 464 on fundamental aspects of forecast system design are occurring at operational centers 465 466 worldwide (Takaya, 2019). One major emphasis consists of methods to represent the 467 uncertainty in initial conditions (bred vector, singular vector, ensemble data assimilation) and model physics (stochastic physics, Leutbecher et al. 2018). In addition, configurations of real-468 469 time forecasts and hindcasts, including ensemble size, ensemble strategy (lagged ensemble 470 with different initial times or burst ensemble with a common initial time) and hindcast period, 471 impact forecast quality and ability to evaluate the performance of the hindcast. Identifying suitable compromises and trade-offs in forecast system design is a challenge under practical 472 473 constraints for operational activities (costs, priorities, timeliness) and demands further research. 474

475

From the modelling perspective, multiple operational centers are moving towards a unified, or "seamless" coupled forecast system that can be applied across timescales from days to seasons or longer. More S2S models are incorporating ocean and sea-ice components, and becoming increasingly complex and complete in representing coupled processes in the Earth system. On the other hand, poor representation of model physics, in particular clouds (Morcrette et al.

Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

2018), results in model drifts and biases in surface land and ocean temperatures, which is a 481 482 long-standing modeling issue that can degrade the skill of S2S predictions (Vitart and Balmaseda, 2017). Improvements in cloud parameterizations (Stan and Straus, 2019) and in 483 484 representing the diurnal cycle of the atmospheric boundary layers are crucial for advancing S2S 485 modeling. The Earth system modeling approach poses another challenge to initialize the ocean and sea ice components with high accuracy; for example there is a relatively large dispersion of 486 487 initialized sea ice fields in current S2S models (Chevallier et al. 2017, Zampieri et al. 2018). 488 Another important S2S modeling issue is predicting the MJO, owing to its importance as a 489 source of subseasonal predictability (H. Kim et al. 2018). Multi-model evaluations have shown that S2S models have difficulties in representing MJO propagation across the Maritime 490 491 Continent. Process-oriented diagnostics (Maloney et al. 2019) have identified a dry bias in the lower troposphere as one of the causes for the poor MJO propagation through weakening the 492 493 horizontal moisture gradient over the Indian Ocean and western Pacific (Lim et al. 2018) and 494 dampening the organization and propagation of the MJO. A recharge process whereby moisture 495 builds up in the lower troposphere during the suppressed convection phase of the MJO, and that is key for MJO propagation around the Maritime Continent in boreal winter, is 496 underrepresented in S2S models due to the dry bias (Kim 2017). Ocean coupling is another 497 498 important process for the MJO (DeMott et al. 2015), and several studies have demonstrated 499 that ocean coupling can improve MJO propagation and enhance predictive skill in models. 500

501 Poor vertical resolution, low model lid height, inadequate orographic and non-orographic
502 gravity wave parameterizations, and biases in the tropospheric mean state (e.g., the location of

stationary Rossby waves) could limit the predictive skill from stratosphere-troposphere 503 504 coupling processes (Tripathi et al. 2015; Butler et al. 2016), but new generations of prediction systems have rapidly improved in many of these areas. Future model development could 505 506 prioritize improved representation of orographic and non-orographic gravity wave drag and an 507 internally-generated QBO (Butchart et al. 2018). Better understanding of stratospheretroposphere coupling processes and the role of the stratosphere on surface skill could be 508 509 gained through case studies and stratospheric nudging experiments (Hansen et al. 2017). 510 Improved observations of the stratosphere (e.g., aerosols and chemistry) and other climate 511 components may improve S2S predictions. Finally, there is potential for modeling of stratospheric ozone chemistry which provides surface temperature predictability on S2S time 512 513 scales due to its influence on high-latitude stratospheric circulation anomalies together with their lagged surface impacts (Stone et al. 2019). Although that may currently be too resource-514 515 intensive due to the many species and reactions that must be modeled, emerging machinelearning techniques may provide pathways for incorporating chemistry-climate information into 516 517 S2S forecasts (Nowack et al. 2018).

518

519 Seasonal to decadal

520 Modeling issues for S2D prediction naturally overlap with those for S2S prediction. However, 521 the longer time scales of S2D prediction lead to a greater emphasis on representing slower 522 climate variations such as ENSO and AMV, and greater attention to reducing model biases in 523 the ocean that may take months to years to develop. Increased model resolution can reduce 524 model biases as illustrated in Fig. 5 (Jia et al. 2015; Müller et al. 2018), and improve skill

(Prodhomme et al. 2016b; Schuster et al. 2019; Infanti and Kirtman 2019), although the greater 525 526 computational cost is not always justified (Scaife et al. 2019). More fundamental strategies involve analyzing/understanding model biases, before attempting to correct them a priori or a 527 posteriori. Such analysis methods include comparing hindcasts with observations and multi-528 529 decadal historical or other simulations to distill causation for model errors, such as in the tropical Pacific (Shonk et al. 2018) or Atlantic (Voldoire et al. 2019). Similarly, errors in modeled 530 variability or teleconnection patterns can be characterized by examining their evolution with 531 532 lead time. Model biases can be corrected both through simple methods such as statistical bias 533 correction and anomaly coupling (Toniazzo and Koseki, 2018), and more complex methods such as supermodeling, through which multiple models exchange information during a climate 534 535 simulation (Shen et al. 2016).

536

537 Performance of S2D predictions is strongly tied to initialization of model components beyond 538 the lower atmosphere. For example, stratospheric initial conditions are a source of seasonal winter NAO skill (e.g., O'Reilly et al. 2019; Nie at al. 2019) as illustrated in Fig. 6, and ocean 539 initial conditions are crucial for skillfully predicting ENSO (Balmaseda and Anderson 2009), as 540 well as decadal variability in the subpolar North Atlantic (Yeager and Robson, 2017; Borchert et 541 542 al. 2018). However, initialization using full-field observational values can lead to initial shocks 543 affecting skill (Kröger et al. 2018) and in such cases initialization combining observed anomalies with the model's own climatology can be beneficial until underlying model errors can be 544 545 reduced (Volpi et al. 2017). Basic initialization strategies continue to be an active research area 546 particularly for decadal prediction (Brune et al. 2018), and methods extending to forecast runs

such as the ensemble dispersion filter which replaces the ensemble members with the 547 548 ensemble mean every three months (Kadlow et al. 2017) are also being explored. Comparisons that apply different initialization methods to the same model can yield valuable insights 549 (Polkova et al. 2019); further issues specific to the initialization of the land, ocean, and sea ice 550 551 components are considered in the next section. 552 553 Tackling these diverse and persistent modeling issues effectively will require sustained effort, as 554 simple model-specific solutions may not cure the underlying problems, and ideally this should 555 involve coordination between the S2S/S2D prediction, climate modelling, and data assimilation communities. 556 557 Initialization issues. 558 559 Atmosphere initialization Accurate atmospheric model initialization is a basic requirement for numerical weather 560 prediction because atmospheric initial conditions are the primary source of predictability on 561 time scales less than a week or two (Fig. 1). It is enabled by sophisticated data assimilation 562 systems that are the result of decades of advancement (Bauer et al. 2015). Subseasonal and 563 564 seasonal prediction systems generally initialize their atmospheric components by such means, 565 with the additional requirement that historical observations must be assimilated similarly to produce reanalyses that are used to initialize hindcasts. Because in situ and remotely sensed 566 atmospheric observations are relatively dense there is generally good agreement between 567

568 different reanalyses for the modern era implying relatively low uncertainty at heights below

about 10 hPa, although temporal inconsistencies can result from changes in observing systems
(Long et al. 2017). Because atmospheric initial conditions contribute less to predicability on
multi-annual time scales, some decadal prediction systems do not initialize the atmosphere
(e.g., Yeager et al. 2018).

573

574 Land initialization

Climatically important land variables such as soil moisture and snow can be initialized by driving 575 576 land surface models with observed atmospheric fields (e.g., Koster et al. 2009; Sospedra-577 Alfonso et al. 2016a) or, more directly, assimilation of land observations principally from satellites (Bilodeau et al. 2016; Muñoz-Sabater et al. 2019; Toure et al. 2018). Yet predictability 578 579 from land surface states is being harvested only to the extent that land initial conditions and the relevant processes are represented realistically in forecast models (Koster et al. 2011; 580 581 Ardilouze et al. 2017). Historically, land surface and atmospheric models are developed separately and their coupled behavior is not calibrated or validated (Dirmeyer et al. 2019), so 582 583 that coupled processes are often not represented accurately (Dirmeyer et al. 2018b). 584 585 There are also observational limitations. In situ measurements of soil moisture are of varying 586 quality and uneven distribution, and are not designed for real-time operational use (Dorigo et 587 al. 2011). Satellite soil moisture monitoring (Entekhabi et al. 2010; Kerr et al. 2010), provides either very shallow or total column measurements including groundwater (Li et al. 2012), and is 588 subject to uncertainties caused by vegetation, etc. (Al-Yaari et al. 2017). By contrast, soil 589 590 moisture in forecast models is mainly a gross reservoir for the surface water balance, and its

variations do not represent all of the observed processes, particularly at sub-grid scales.

592 Therefore model soil moisture is only a crude representation of reality, although it still contains

593 useful information that can be largely consistent across different land models (Koster et al.

594 2009).

595

596 Climate forecasts can be improved by making high-quality land state observations an 597 operational priority for real-time reporting, and planning for long-term continuity in satellite 598 monitoring (Balsamo et al. 2018). This includes vegetation, especially as its interannual 599 variability and cycles of agricultural planting and harvest are not represented and can affect surface fluxes and predictions (Alessandri et al. 2017). In addition, realistic snow initialization 600 601 can positively impact subseasonal predictions of surface temperatures (e.g., F. Li et al. 2019). Along with coupled land-atmosphere model development (Santanello et al. 2018), such efforts 602 603 would facilitate improved predictions on weather to subseasonal time scales, as demonstrated by numerous forecast model-based sensitivity studies such as that of Koster et al. (2011). 604 605 Ocean and sea ice initialization 606 The importance of initializing the oceans stems from their relatively long thermal and dynamical 607 608 time scales, through which they play an essential role in S2D climate predictability (Cassou et al. 609 2017). In addition, the oceans can influence S2S variability, for example through air-sea interactions affecting the MJO (DeMott et al. 2015) and mesoscale eddy impacts on 610 atmospheric circulation (Saravanan and Chang 2019). Predicting future ocean evolution, 611 612 especially on S2D time scales, requires estimates of 3D ocean states for initialization. This in

turn requires a data assimilation method (usually in conjunction with a dynamical model) to 613 614 constrain ocean state estimates based on available observations. Similar considerations apply to state estimates of sea ice. Comparisons of different ocean and sea ice state estimates as in 615 Fig. 7 can point to variables and regions for which they are most robust, as well as to where 616 617 uncertainties are relatively large (Balmaseda et al. 2015; Chevallier et al. 2017). Observing system experiments in which certain observations are withheld have shown for example that 618 data from tropical ocean moorings positively impacts state estimates even when Argo float 619 620 data is also available (Fujii et al. 2015).

621

622 Recent enhancements in observing capabilities are enabling improvements in ocean and sea ice 623 state estimates, potentially leading to more accurate initial conditions and hence better forecasts. For example, assimilation of satellite measurements of sea surface salinity (SSS) leads 624 625 to improvements in tropical Pacific ocean states and ENSO forecasts in experiments using an intermediate-complexity coupled model (Hackert et al. 2019), whereas assimilation of satellite-626 627 derived sea ice thickness (SIT) measurements has shown potential for improving sea ice forecasts in operational seasonal forecasting systems (Chen et al. 2017; Blockley and Peterson, 628 2018). A major limitation is that these data sources have been available for less than a decade, 629 630 whereas considerably longer hindcast periods are required for forecast post-processing and skill 631 assessment, and temporal consistency of observational data used for initialization is required to avoid artificial biases between hindcasts and forecasts. Forecasts thus continue to be initialized 632 typically without assimilation of SSS or SIT, from initial conditions that deviate appreciably from 633

634	available observations especially for SIT (Uotila et al. 2019). This motivates alternative
635	approaches for initializing SIT over multidecadal hindcast periods (Dirkson et al. 2017).
636	

637 Coupled data assimilation

638 The atmosphere, land, ocean and sea ice components of climate prediction models have often been initialized individually, without coupling. However, such an approach does not make 639 optimal use of observations, which may exert influences across the interfaces of the model 640 641 components. In addition, physical inconsistencies between the separately initialized 642 components may lead to rapid adjustments, or shocks. To overcome these limitations attention has increasingly turned toward developing coupled data assimilation methods that treat 643 644 multiple components, such as atmosphere and ocean, simultaneously using observations from each (Penny and Hamill 2017). Such methods are termed weakly or strongly coupled (Penny et 645 646 al. 2017). Weakly coupled methods apply assimilation independently to each model component within the coupled model, so that the components may exchange information across their 647 648 interfaces. Such techniques have shown promise for reducing shocks (Mulholland et al. 2015), and have begun to be applied operationally (e.g., Browne et al. 2019). Strongly coupled 649 methods apply assimilation to multiple model components in an integrated manner, so that 650 651 observations assimilated in one component can directly influence others. Such methods 652 remain experimental and thus far have been applied mainly in simplified models (e.g., Penny et al. 2019). 653

654

655 **Ensemble predictions and forecast information.**

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Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

657 In contrast to ensemble weather forecasts, a consolidated verification strategy for S2S predictions is not yet established, and developing such a framework that encompasses 658 659 important forecast attributes such as accuracy, association, discrimination, reliability, and 660 resolution has thus emerged as a priority (Coelho et al. 2018). (Accuracy measures error, or distance between forecast and observed values; association measures strength of the linear 661 relationship between forecast and observation as through temporal or spatial correlations; 662 663 discrimination measures by how much forecasts differ given different outcomes; reliability 664 measures how well forecast probabilities correspond to observed frequencies of occurrence; resolution measures by how much outcomes differ given different forecast probabilities. 665 666 Forecast quality encompasses all these attributes, whereas skill indicates quality relative to some benchmark such as persisted anomalies or climatological probabilities.) As for seasonal 667 668 predictions, a purpose of S2S hindcasts is to provide a larger sample for more confident 669 verification statistics than real time forecasts because they cover more years. However, 670 because S2S hindcasts are initialized using re-analysis and most often have a smaller ensemble size, their verification generally underestimates real-time forecast quality. Operational centres 671 are encouraged to compute and monitor verification statistics based both on hindcasts and 672 673 real-time forecasts.

674

As has been demonstrated for seasonal prediction, S2S multi-model ensembles (MMEs)
generally outperform individual models (Vigaud et al. 2017; Pegion et al. 2019). Currently, the

S2S and SubX MME projects are providing testbeds for research³ as well as a foundation for 677 678 operational use (Vitart and Robertson 2019; Pegion et al. 2019). One focus for exploiting such datasets is developing calibration procedures, post-processing steps that improve the 679 properties of probabilistic forecasts, to enable S2S ensemble forecasts to provide reliable 680 681 probabilities for particular conditions occurring or thresholds being exceeded, especially for extreme events. The varied current choices among S2S project modelling systems for hindcast 682 and near real time initialization dates, hindcast period and ensemble size is, however, limiting 683 684 advances in developing multi-model calibration and combination procedures. In addition, the value of these datasets for research would be enhanced if more comprehensive stratospheric 685 data were to be available across models. 686

687

S2S ensemble forecasts have shown promise in providing useful predictions and early warnings 688 689 for high impact climate and weather events including severe heat waves and cold spells, as well as regional probabilities of the occurrence of tropical storms as illustrated in Fig. 8 (Vitart and 690 691 Robertson 2018). Examples include severe cold conditions over Europe associated with the negative phase of the NAO, whose useful predictability into week 3 is enhanced by tropical-692 extratropical teleconnections resulting from MJO activity (Ferranti et al. 2018), and atmospheric 693 rivers, plumes of intense water vapor transport that often trigger weather and hydrologic 694 695 extremes and are especially predictable at lead times of 3 to 5 weeks during certain MJO and QBO phase combinations (Baggett et al. 2017). While modest overall skill at ranges longer than 696

³ Hindcast and near real-time forecast data are available from S2S at <u>www.s2sprediction.net</u> and from SubX at http://iridl.ldeo.columbia.edu/SOURCES/.Models/.SubX/.

a week has been found for S2S predictions of springtime Sahelian heat waves including

698 measures of heat stress, such conditions following a strong El Nino were accurately forecast,

pointing to the tropical Pacific as a source of predictability for extremes in that region (Batté etal. 2018).

701

A global precipitation hindcast quality assessment of the S2S prediction project models (Fig. 9) 702 703 was performed by de Andrade et al. (2019). Sub-seasonal prediction quality is modulated by 704 the MJO, QBO, ENSO in the tropics, changes in large-scale flow in the extra-tropics and 705 stratospheric tropical and extratropical variability (Butler et al. 2019). It is therefore important 706 to estimate the predictive skill of such events and identify their impacts on predictions of 707 weather and weather extremes. Evaluating the conditional prediction quality associated with 708 the key low frequency variability modes is instrumental for better understanding S2S 709 predictability mechanisms. For example, MJO predictive skill in the S2S MME ranges between 12 to 36 days and is affected both by the MJO amplitude and phase errors (Vitart 2017; Lim et 710 711 al. 2018; H. Kim et al. 2018). Communicating these variations in forecast quality, including if the forecasts are no better than climatology, is extremely important as users with such knowledge 712 can better utilize and benefit from the forecast information. Furthermore, capitalizing on 713 714 "windows of opportunity" when skill is especially high increases the value of S2S forecasts 715 (Mariotti et al. 2020), and motivates their frequent initialization (ideally daily).

716

717 Seasonal to decadal

Limited forecast quality in current S2D ensemble prediction systems motivates research initiatives that focus on extracting skillful and reliable information from the large amounts of forecast and hindcast data available to potential users⁴.

721

722 One emerging theme of such research is that S2D prediction systems sometimes underestimate 723 the predictable signal (Eade et al. 2014; Scaife and Smith 2018). As a result, very large ensembles that effectively filter out unpredictable noise demonstrat higher skill in predicting phenomena 724 725 such as the winter NAO (Scaife et al. 2014b; Dunstone et al. 2016) and seasonal to multi-annual regional precipitation variations (Dunstone et al. 2018; Yeager et al. 2018) than was previously 726 thought possible. While very large ensemble sizes hold value for isolating weak predictable 727 728 signals, much smaller ensemble sizes are sufficient for skillful prediction of tropical SST, for which signal to noise ratios are much larger (Zhu et al. 2015). The causes of unrealistically low modeled 729 730 predictable signals (sometimes called the "signal to noise paradox") remain under investigation. Two hypotheses stemming from hindcast experiments are that winter NAO skill is enhanced by 731 732 skillful prediction of a QBO teleconnection that is too weak in models (O'Reilly et al. 2019), and that transient eddy feedbacks are too weak in models (Scaife at al. 2019). Others based on simple 733 models suggest that the NAO predictable signal is too weak because climate models switch 734

⁴ Seasonal hindcast data from the WCRP Climate-system Historical Forecast Project (CHFP; Tompkins et al. 2017) are available at http://chfps.cima.fcen.uba.ar/access.php, and from the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) including real-time forecasts at https://iridl.ldeo.columbia.edu/SOURCES/.Models/.NMME/. Decadal hindcast data from the WCRP Coupled Model

Intercomparison Project Phases 5 and 6 are available via https://esgf-node.llnl.gov/projects/cmip5/ and https://esgf-node.llnl.gov/projects/cmip6/.

between NAO regimes too rapidly (Strommen and Palmer 2019), or exhibit less persistent NAO
variability than is observed (Zhang and Kirtman 2019).

737

In the case of the winter NAO which is a key source of variability over the mid-latitude North Atlantic and Europe, another approach to extract relevant information from over-dispersive ensembles that leads to improved skill is to subsample ensemble members that are close to a "first guess" statistical prediction of the NAO (Dobrynin et al. 2018); subsampling has shown potential for improving European summer forecasts as well (Neddermann et al. 2019).

743

Estimating and realizing the predictability of key modes of variability is still a major challenge at 744 745 S2D time scales. ENSO is considered one of the most predictable phenomena on multi-seasonal time scales, but longer-range skill has been viewed as limited. However, multi-year ensemble 746 747 predictions have shown evidence of skill in predicting long-lasting La Niña events that follow warm events up to 24 months ahead (DiNezio et al. 2017; Luo et al. 2017). Challenges in the 748 749 initialization of such longer time scale predictions remain, as evidenced by multi-year predictions in which skill for SST and precipitation over land improves with lead time in some areas, 750 suggesting that short-term adjustments following initialization may tend to degrade skill (Yeager 751 752 et al. 2018).

753

Calibration of ensemble forecasts is a necessary step to reduce the areas for which S2D forecasts
are unreliable and potentially misleading. Combinations of several forecasting systems such as
the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) are now routinely used

to increase ensemble reliability and improve forecast skill. Several recent efforts have explored
weighted multi-model calibration methods to combine ensembles from different models in order
to improve probabilistic seasonal forecasts for temperature and precipitation anomalies as well
as forecast of extremes (Becker 2017). Calibration methods have also been developed for
ensemble decadal hindcasts to adjust both the bias and ensemble spread with a parametric
dependency on lead time and initialization time (Pasternack et al. 2018). Such methods are found
to improve both the conditional bias and probabilistic skill of decadal hindcasts.

764

765 Climate forecasts for decision making.

766 Subseasonal to Seasonal

767 Many decisions are made on the S2S forecasting timescale, which sits between weather forecasts and S2D climate outlooks; therefore the continued development of S2S forecasts has 768 769 the potential to benefit many sectors of society (Fig. 10). S2S forecasting is a rapidly maturing discipline, with emerging recognition for both the need and the potential use of forecasts on 770 this timescale (White et al. 2017). While S2S forecasts are increasingly being used in 771 government as well as a range of sectors including agriculture, energy, finance, health and 772 water resource management – more engagement between S2S forecasters and end users is 773 774 needed to increase the wider awareness and uptake of S2S forecasts. 775 Although scientific knowledge gaps, computational capacity, and gaps in observations and 776 modeling currently limit the use of S2S forecasts to some degree, by increasingly placing 777

decision makers at the forefront of S2S forecast development, an improved dialogue between

Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

S2S forecasters, developers and end users will accelerate the awareness and application of S2S
forecasts to real-world decision-making.

781

782	To support the increased use of S2S forecasts for decision-making, the following
783	recommendations were identified for action following the Boulder conference:
784	• A summary of existing stakeholder case studies is planned to be created to demonstrate past
785	and ongoing 'success stories', and support better engagement with end users and
786	stakeholders. As the S2S forecast needs and associated performance varies greatly between
787	different sectors and users, the wider community is increasingly working together on the co-
788	design and production of S2S predictions in order to better meet user needs. Several
789	applications of S2S forecasts are now being developed in different disciplines, such as the
790	EU-funded <u>S2S4E</u> project in the energy sector, a quasi-operational excess heat outlook
791	system in the health sector (Lowe et al. 2016), and S2S hydrologic prediction in the water
792	management sector. These efforts need to be catalogued and disseminated to guide further
793	user-driven decision-support products, and to support the continued development of S2S
794	forecast, verification metrics and related services.
795	• Systematically assessing the relative skill (or lack thereof) of forecasting a series of historical
796	high-impact events, such as heat waves, extreme rainfall events, or wildfires, on the S2S
797	timescale would be a useful way to help demonstrate the potential of S2S forecasts to
798	decision-makers across multiple sectors. At present, such case studies are often ad-hoc and
799	typically not published in the wider literature; however, a collaborative effort that brings
800	together a set of demonstrable case studies, involving both forecasters and end users, would

fill this gap. For example, a series of 'tailored narratives', or 'storylines' (approaches that
construct stories of plausible, non-probabilistic climatic futures that relate to a specific
person or sector to counter perceived barriers; e.g., Hazeleger *et al.* 2015), may aid in the
understanding of what S2S forecasts may deliver in the future.

To support the co-design, uptake and use of S2S forecasts, S2Sapp.net is currently being
 established as a new network of researchers, modellers and practitioners – an 'open to all'
 global community with a shared aim of exploring and promoting cross-sectoral services and
 applications of this new generation of forecasts from across government, academia, and the
 private sector.

810

811 Seasonal to decadal

Research efforts are assessing the value of S2D forecast information for many applications, and
initiatives such as the WMO's Global Seasonal Climate Update⁵ and Annual to Decadal Climate
Update (Kushnir et al. 2019) are making such information more widely available. However,
consultation with decision makers is essential in order to tailor forecast information to the needs
and expectations of users.

817

Fisheries management is one application for which S2D forecast information holds promise (Tommasi et al. 2017). This is due to reasonable skill for ocean prediction in regions of interest, coupled with strong influences of S2D climate variability on fish populations. Case studies

⁵ https://public.wmo.int/en/our-mandate/climate/global-seasonal-climate-update

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employing fisheries management decision frameworks have shown that SST forecast information 821 822 can potentially increase fishery yields while reducing the risk of population collapse from combined effects of environmental factors and overfishing. However, significant challenges 823 remain for fully realizing this potential. These include a need for improved initialization and 824 825 reduced model errors in key ocean regions such as the US Northeast continental shelf, dynamical downscaling in cases where important spatial scales are not resolved by global models, and 826 827 sufficiently accurate observational data for hindcast verification on these scales. In addition, 828 incorporating biogeochemistry and marine ecosystem components into prediction systems will 829 expand their potential capabilities, while posing additional verification challenges.

830

831 Another current focus of application-oriented research is water management. Global climate prediction models have been shown to have skill in predicting the next winter season's snowpack 832 833 throughout much of the western US, where spring snowmelt is an essential water resource (Kapnick et al. 2018; Sospedra-Alfonso et al. 2016b). Because temperature influences snowmelt 834 835 and runoff efficiency, skill in seasonal temperature forecasts can provide improved skill for seasonal water supply forecasts in this region (Lehner et al. 2017). Seasonal forecast skill has also 836 been demonstrated for monsoon rainfall (e.g., Jain et al. 2019) and drought (Hao et al. 2018) with 837 potential to inform water management decisions in many regions of the globe. Decadal forecasts 838 839 potentially can meet planning horizon needs but currently are less familiar to water managers than seasonal forecasts and long-term climate projections. Efforts to apply decadal climate 840 information for water management decisions have included assessing the role of decadal modes 841 842 of variability, and using statistically downscaled decadal predictions as hydrological model inputs.

B43 Developing information that is credible and compatible with existing decision frameworks is an
important consideration (Towler et al. 2018).

845

Additional sectors for which S2D forecasts are being assessed for decision making include agriculture (Klemm and McPherson, 2017), energy (demand & wind power generation, Clark et al. 2017; Lledó et al. 2019), tropical cyclone (Bergman et al. 2019) and coastal flooding (Widlansky et al. 2017) preparedness, Arctic marine transportation (Stephenson and Pincus 2018), wildfire risk (Turco et al. 2019), and food security (Funk et al. 2019).

851

Initiatives to develop and deliver climate forecast information range in scale from international, 852 853 regional and national (e.g., Marotzke et al. 2016), to individual users, all of which aim to provide forecast information having practical value for decision makers. In all cases, it is crucially 854 855 important that uncertainties are adequately quantified and conveyed in order to avoid any false sense of certainty and to build trust in forecast information providers, although sometimes this 856 requires overcoming a preference of users for deterministic information. Additional 857 considerations are that expectations of users need to be conditioned to generally modest levels 858 of skill, but that this information can nonetheless be advantageous when applied consistently in 859 860 the long term. The likelihood that climate forecast information gets used increases when efforts 861 are made to build relationships with potential users, and dialogs are opened to enable forecast products to be co-designed (Kolstad et al. 2019). 862

863

864 **Cross-cutting issues in S2S and S2D prediction.**

865 Initialization shock and model error

Model biases are an endemic modeling issue that is common across S2S and S2D prediction time scales and influence all aspects of the prediction systems – complicating ingestion of assimilated observations, degrading skill, and necessitating post-processing steps such as bias correction and calibration for product development and delivery.

870

Model biases begin to develop on fast time scales and lead to drifts from the climate 871 872 represented by the initial conditions to that of a model's biased equilibrium state. It has been 873 extremely hard to understand the mechanisms behind these drifts, and further, pathways for their diagnosis are not clear although some progress is being made (Sanchez-Gomez et al. 2015; 874 875 Shonk et al. 2018; Voldoire et al. 2019). Such difficulties arise due to non-linear interaction between various physical processes that are parameterized, and because biases could be non-876 877 local in their origin. Long time scales before models' equilibrium states are attained make understanding the causes of drifts even harder. The Boulder meeting recognized that the 878 879 S2S/S2D prediction community needs to pay particular attention to developing pathways for understanding the onset of model biases and put together mechanisms (such as summer 880 schools) to train the next generation of scientists with interest and expertise in climate 881 modeling and model diagnostics. 882

883

Initialization shocks that arise from imbalances in initial states with respect to the formulation
of the model and can be caused by limitations of observations and data assimilation as well as
model biases were also recognized as a major issue, particularly in the context of decadal

predictions. Initialization shocks result in the degradation of initial information that may be the 887 888 primary source of predictability for the subsequent forecast. Even after considerable research and investment in decadal predictions it is still not clear what may be best approaches, such as 889 890 between full field vs. anomaly initialization, to retain predictive information in the initial state 891 while minimizing the influence of initial shocks on the subsequent forecast. The continuing prominence of model drift and initial shocks as important issues reinforces a long held 892 sentiment that these are outstanding problems that need to be studied more systematically if 893 894 progress in translating inherent predictability into prediction skill is to be made.

895

896 S2S and S2D research interactions

897 The examples of interaction among modes of variability across S2S and S2D time scales noted earlier emphasize the fact that continued interaction and communication across the S2S and 898 899 S2D research communities will be important to make progress. Furthering our understanding of time-scale interactions will require investments in process level understanding of these 900 phenomena and will not only benefit our understanding about their lower-frequency variations 901 but will also contribute to improved process level diagnostics of model simulations. Better 902 understanding of time-scale interactions is likely to require the use of a hierarchy of models, 903 904 such as simple linear models to investigate the characteristics of tropical-extratropical 905 interactions (including their influence on storm tracks), to diagnose possible causes for errors in their representation in complex GCMs (Dias et al. 2019). 906

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Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

37

908	Another aspect of research interactions across time scales is quantifying the fidelity of models
909	in S2S and S2D prediction as well as projections of climate on longer time scales based on their
910	simulation and prediction of shorter time-scale phenomena. The advantage of such an
911	approach is that much larger samples for predictions of shorter time-scale phenomena are
912	available, and an assessment of the reliability of such predictions can be used to build
913	confidence in prediction on longer time-scales. Theoretical basis for extrapolating the reliability
914	of forecasts across different time scales may also require the use of a hierarchy of models
915	(Weisheimer and Palmer 2014; Christensen and Berner 2019).
916	
917	Research and operations
918	Post-processing to improve forecast quality is an important area of research that bears directly
919	on operational activities. Post-processing is necessary because biases in forecasts can be as
920	large as the predicted signal, and therefore require the use of bias correction and calibration
921	techniques to adjust real-time predictions before their delivery to the users. These
922	requirements are shared across sub-seasonal to decadal prediction time-scales, however
923	because of different levels of experience (seasonal predictions having the longest history) the
924	opportunity for cross-community interactions was recognized. Some aspects for post-
925	processing are specific to time-scale, for example, bias correction for decadal predictions
926	requires methods to account for the non-stationarity of climate, and research needs to develop
927	such methods were stressed.
928	

38

Necessity for post-processing requires an extensive set of hindcasts to accompany real-time predictions. Because of limited resources, decisions about hindcast period, ensemble size and forecast start dates are not straightforward and call for further guidance from the research community. Such questions about the operational infrastructure for long-range prediction systems, including ensemble generation techniques and recommendations for harmonizing hindcast and real time forecast production, provide an opportunity to link operational and research communities that was highlighted during the conference.

936

937 Product development and communicating forecasts to the user community is also a common thread across the S2S and S2D communities. Communication of probabilistic forecast 938 information (including confidence in the forecast based on past verifications) to users for their 939 940 decision making has been a challenge, and once again there is much to be gained from lessons 941 learned from the experiences of different communities. Similar challenges and opportunities also exist in the context of product development that incorporate user needs based on an 942 ongoing dialog from the very start of the process. In addition, users often wish to have 943 information on smaller spatial scales than are represented in global climate models. For such 944 applications either statistical or dynamical downscaling is possible and can be effective in 945 reducing local climatological biases, although clear demonstrations that downscaling can 946 947 improve the skill of climate predictions remain elusive (e.g., Manzanas et al. 2018).

948

949 In summary, research needs for further development of operational infrastructure, product

950 generation and communication of probabilistic forecasts were themes often repeated during

951 the conference.

952

953 Conclusions and the future of subseasonal to decadal prediction

This paper has outlined many commonalities in the prediction of weather and climate across time scales and Earth system components, and through the value cycle from basic research to operational delivery.

957

The Earth's weather and climate is inherently chaotic and challenges the best currently 958 959 available modeling capabilities. There remains however untapped skill, and realizing this skill 960 will require improvements on numerous fronts. These include fundamental understanding of 961 fine-scale processes, leading toward their robust parameterization; accurately representing property exchanges across Earth system components through realistic coupling limiting 962 systematic errors; sustained Earth observing systems and advanced data assimilation methods 963 enabling balanced initial conditions that avoid shocks and mitigate model drifts; and innovative 964 numerical and ensemble generation techniques to address model scalability issues. Additional 965 966 important avenues toward improved services include development of probabilistic information 967 for high impact weather and climate events including unprecedented extremes, and optimal post-processing and data fusion to add value to multi-model ensembles, among many others. 968 969

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 973 974 The joint WWRP-WCRP conferences in Boulder clearly demonstrated the benefit in bringing 975 relevant stakeholders together to cross-fertilize their experience, knowledge, respective issues 976 and working cultures, which will surely help frame a new and vibrant research portfolio, and 977 inspire the next generation of science leaders to ensure that society has access to the best 978 possible weather and climate prediction science. 979 980 ACKNOWLEDGEMENTS 981 The International Conferences on Subseasonal to Decadal Prediction on which this paper is 982 based were sponsored by: US CLIVAR, NSF, UCAR, NCAR and its Climate and Global Dynamics 983 Laboratory (CGD), NOAA's Climate Variability and Predictability (CVP) and Modeling, Analysis, 984 Predictions and Projections (MAPP) Programs, Copernicus Climate Change Service, IPSL, and 985 WWRP/WCRP's Subseasonal-to-Seasonal (S2S) Prediction Project.
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985 WWRP/WCRP's Subseasonal-to-Seasonal (S2S) Prediction Project.
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These challenges are broad but so are opportunities for steady progress, ranging from curiosity-

driven science to the systematic model evaluation and improvement in a collaborative and

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1693 **SIDEBAR 1:**

1694 Hindcast and forecast quality assessment (or, "the unexamined life is not worth living"). Besides helping to inform decision making, the careful assessment of forecast quality is critical 1695 to guiding forecasting improvements, but has many and varied considerations. Simply 1696 1697 answering the question "is this forecast better than that one?" is complicated, as the appropriate skill metric or means for comparison is not always obvious. Some recent studies 1698 1699 have focused on newer statistical methods for comparing one forecast to another. One 1700 relatively simple approach is the random walk test (DelSole and Tippett 2016), illustrated in Fig. 1701 SB1. This method is applicable to a wide range of measures such as a score based on the earth mover's distance metric (Düsterhus 2019), while also being robust and fair in its discrimination. 1702 1703 1704 The utility of forecast assessment can be illustrated through two very different applications of 1705 seasonal forecasts: sea-ice and hurricanes. The assessment of seasonal sea ice forecasts is complicated by the low quality of sea-ice observations, but nevertheless reveals that initializing 1706 1707 sea-ice thickness using observational data sets generally improves the prediction of Arctic sea-1708 ice extent and edges (Blockley et al. 2018). Comparison of multi-annual forecasts of Atlantic hurricane activity obtained by direct tracking of storms in decadal hindcasts and through a 1709 1710 hybrid approach combining predicted SSTs and observed statistical relations finds that each 1711 approach is skillful, especially hybrid forecasts based on a SST index for AMV (Caron et al. 2018). 1712

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Accepted for publication in Bulletin of the American Meteorological Society. DOI 10.1175/BAMS-D-19-0037.1.

A robust assessment of model performance should include the model's simulation of climate 1714 1715 modes and teleconnection patterns such as ENSO, MJO and NAO, since they are major sources of predictability and errors representating their patterns or strength (e.g., Yang and DelSole 1716 1717 2012; Vitart 2017) can degrade skill in affected regions (Gleixner et al. 2017; Lu et al. 2017). In 1718 cases where modeled teleconnection patterns are imperfect, forecast skill may be improved by means of statistical methods that use model forecasts of relevant climate modes such as ENSO 1719 1720 as predictors (e.g., Strazzo et al. 2019). It remains desirable, however, for models to improve so 1721 that their simulated teleconnection patterns are sufficiently realistic that such corrections are 1722 not needed.

1723

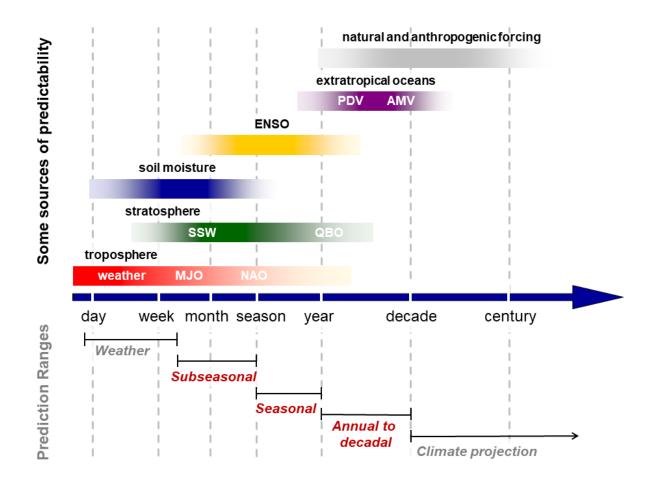
1724 **SIDEBAR 2:**

1725 Frontiers in Earth system prediction.

1726 New frontiers in S2D prediction have been enabled by Earth system models (ESMs, Flato 2011) that represent the carbon and other biogeochemical cycles in addition to the physical climate 1727 1728 system. These frontiers include prediction of ocean and land carbon sinks and biogeochemistry and their important contribution to global carbon storage, as well as ecosystem services. The 1729 world's oceans are a fundamental regulator of global carbon storage and variability. The 1730 1731 strength of ocean carbon uptake, together with uptake of carbon by the land, determines the 1732 fraction of anthropogenic emissions remaining in the atmosphere, and hence modulates present and future warming. Observed global mean ocean carbon uptake shows variability on 1733 1734 decadal time scales that can be represented by ESMs in which physical climate variables are 1735 assimilated (H. Li et al. 2019).

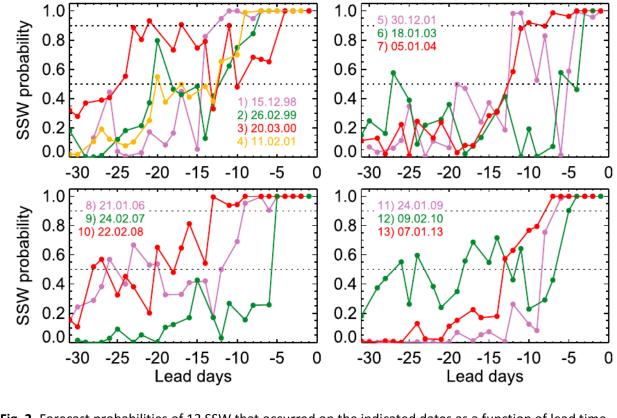
1737	ESM simulations indicate that internal variability of the ocean carbon uptake on decadal
1738	timescales is as large as the forced climate change trend (Li and Ilyina 2018), pointing to the
1739	potential importance and utility of decadal carbon cycle predictions. Decadal predictions from a
1740	number of ESMs are assessing the predictability of the ocean and land carbon sinks and other
1741	ocean tracers such as dissolved oxygen. These forecasts are part of the Decadal Climate
1742	Prediction Project (Boer et al. 2016) and international programs such as the World Climate
1743	Research Programme's Grand Challenge on Carbon Feedbacks (Ilyina and Friedlingstein 2016).
1744	Initial results based on individual models have demonstrated potential predictive skill, assessed
1745	through verification against the assimilating reconstructions that provide initial conditions, for
1746	ocean carbon uptake in certain regions such as the North Atlantic reaching up to 7 or more
1747	years (Li et al. 2016; Lovenduski et al. 2019), and skill in predicting actual variations estimated
1748	from observations (Fig. SB2) has been demonstrated (Li et al. 2019).
1749	ESM-based studies also point to the drivers of this predictability. Air-sea CO_2 flux mainly varies
1750	due to pCO_2 changes in the ocean. While thermal influences on pCO_2 play a role in shorter term
1751	predictability, predictability beyond 3 years is maintained mainly by nonthermal influences of
1752	ocean circulation and biological modification of surface dissolved inorganic carbon and
1753	alkalinity (Li et al. 2019; Lovenduski et al. 2019).
1754	
1755	Investigations in progress are finding potential for multi-annual prediction of additional
1756	biogeochemical fields such as net primary productivity and interior dissolved oxygen
1757	concentrations. In addition, potential predictability and skill for terrestrial carbon uptake have

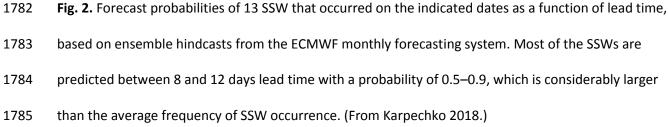
1758	also begun to be assessed, with promising initial results (N. Lovenduski 2019, personal
1759	communication). These examples demonstrate that skillful multi-year prediction is likely
1760	achievable for biogeochemical and ecological Earth system components, and open prospects
1761	for the utilization of such information although significant challenges including the paucity of
1762	long term observational data for initialization and verification will need to be overcome.
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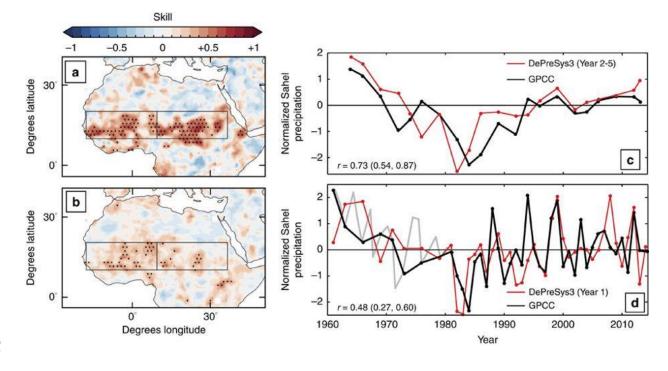


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Fig. 1. Schematic depiction of temporal ranges and sources of predictability for weather and climate 1771 1772 prediction. The subseasonal range encompasses the S2S time scales, and the seasonal and annual-todecadal ranges the S2D time scales, that are considered in this paper. Longer multi-decadal and 1773 1774 centennial ranges derive predictability mainly from forcing scenarios rather than initial conditions, and 1775 are typically represented through climate projections originating from historical simulations begun in 1776 preindustrial times rather than predictions initialized from more recent observation-based climate 1777 states. Some important sources of predictability and approximate time scales over which they are most 1778 influential on surface climate are indicated in the upper portion of the figure; acronyms are defined and 1779 associated phenomena are discussed in the main text.

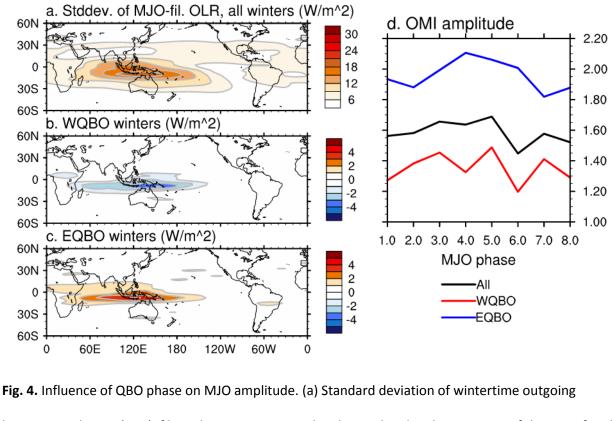






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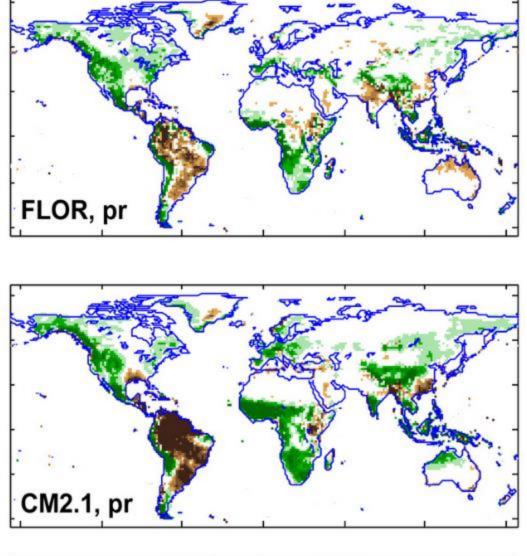
Fig. 3. Skill for predicting linearly detrended Sahel summer rainfall in years 2-5 (upper panels) and year 1
(lower panels) in DePreSys hindcasts. Panels (a)-(b) show spatial distributions of anomaly correlation
coefficients with stippling indicating 95% significance. Panels (c)-(d) show time series of normalized
predicted and GPCC observed rainfall anomalies in the Sahel region outlined by the boxes in the maps,
with correlations and their 5–95% confidence intervals indicated. (From Sheen et al. 2017.)



longwave radiation (OLR), filtered to retain temporal and spatial scales characteristic of the MJO, for all
winters in 1979 to 2012. Differences from these values in winters characterized by QBO westerly
(WQBO) and easterly (EQBO) phases are shown (b) and (c) respectively. (d) Amplitude of an OLR-based

1795

1799 MJO index (OMI) as a function of MJO phase for WQBO, EQBO and all winters. (From Yoo and Son 2016.)



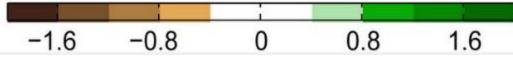


Fig. 5. Impact of resolution on precipitation biases in GFDL seasonal prediction models. Atmospheric
resolution is approximately 50 km with 32 levels in FLOR (upper panel), and approximately 200 km with
24 levels in CM2.1 (lower panel), whereas ocean resolution is approximately 100 km in both models.
Higher atmospheric resolution in FLOR reduces precipitation biases in numerous regions including much
of the tropics. Annual mean biases over land in mm day⁻¹ based on 1981-2010 CMAP observations are
shown. (After Jia et al. 2015.)

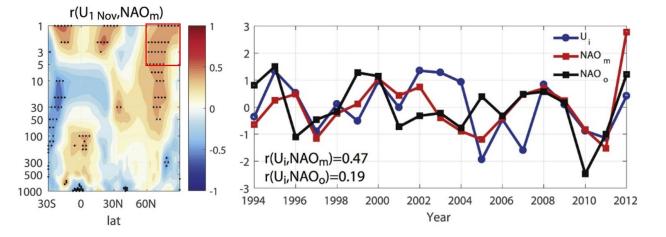
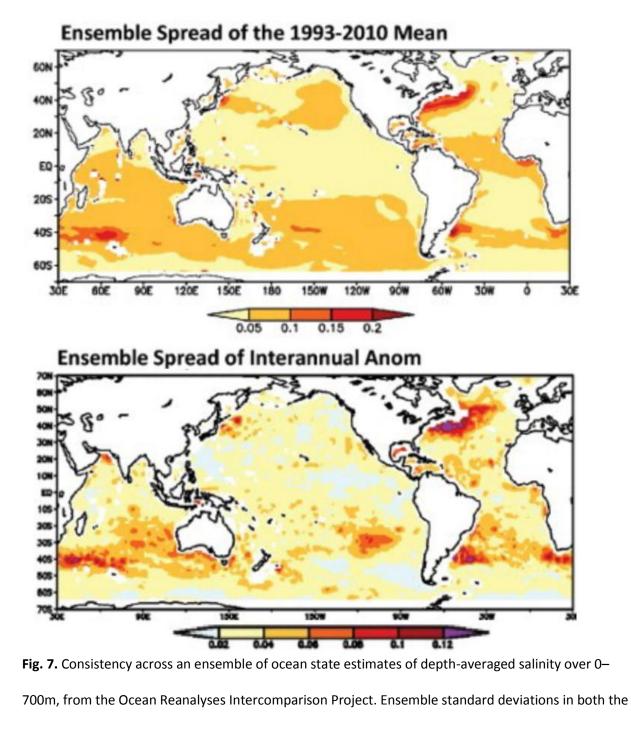


Fig. 6 Connection between stratospheric initial conditions and predicted winter NAO for UK Met Office 1809 1810 GloSea5 predictions initialized 1 November 1995-2012. Left: correlation between initial zonal wind 1811 anomaly on 1 November and ensemble mean model-predicted surface NAO index (NAO_m) during DJF. 1812 Black dots represent values significant at $\alpha = 0.05$ confidence based on one-tailed test, and mean values 1813 within the red box define an index U_i. Right: Annual standardized U_i (blue), NAO_m (red) and observed 1814 surface NAO index, NAO_o (black). The correlation of U_i with NAO_m, indicated at lower left, is significant at 1815 α = 0.05 confidence whereas the lower correlation of U_i with NAO_o is not unexpected based on signal to 1816 noise considerations and that there is only one realization of observations. The larger correlation of 1817 predicted and observed winter NAO values $r(NAO_m, NAO_o)=0.62$ suggests that additional sources of 1818 predictability exist. (After Nie et al. 2019.)



1822 1993-2010 means (upper panel) and interannually varying monthly anomalies (lower panel) tend to be

- 1823 largest in eddy active regions such as the Gulf Stream, and less well-observed regions such as the
- 1824 Southern Ocean. These differences across state estimates are reflective of uncertainties in ocean initial
- 1825 conditions. (After Balmaseda et al. 2015.)

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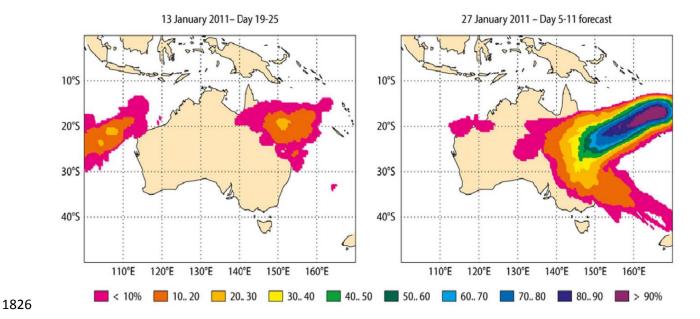
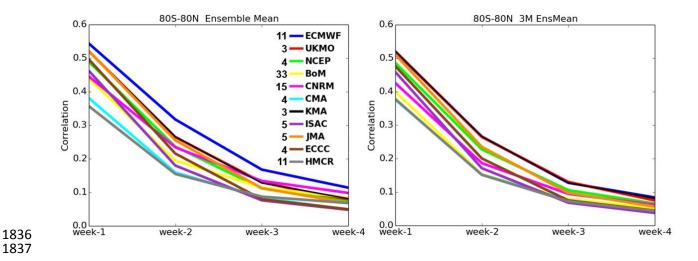


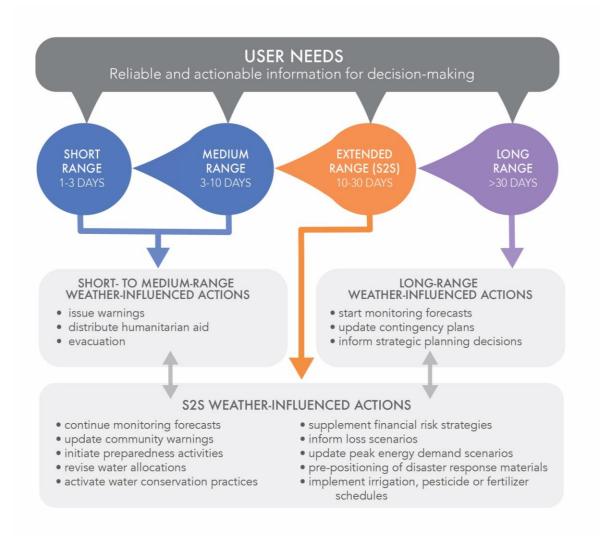
Fig. 8. Elevated probabilities of tropical cyclone occurrence during 31 January to 6 February 2011, based
on ECMWF ensemble forecasts forecast starting 13 January with 18 day lead time (left), and 27 January
with 4 day lead time (right). Destructive Cyclone Yasi made landfall in northeastern Australia on 3

- 1830 February 2011 as a destructive category 5 storm. (Adapted from Vitart and Robertson 2018).
- 1831

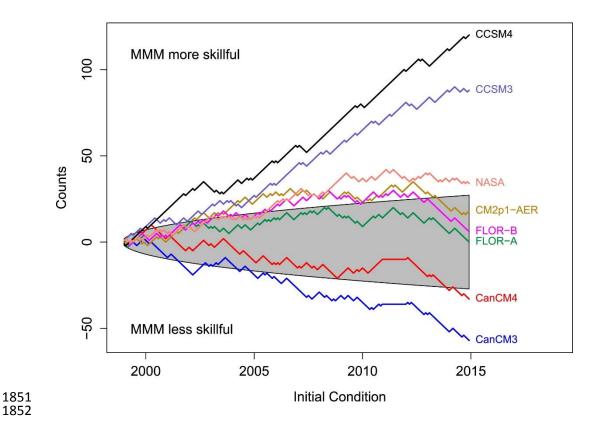
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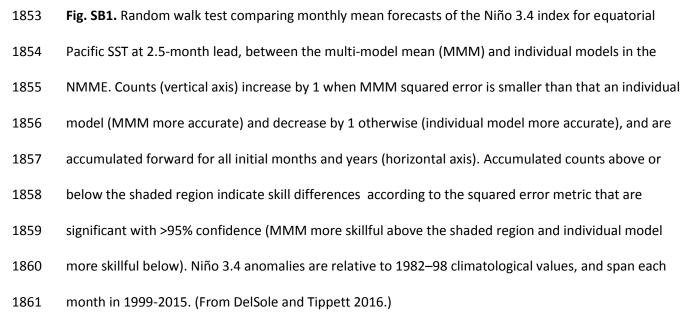


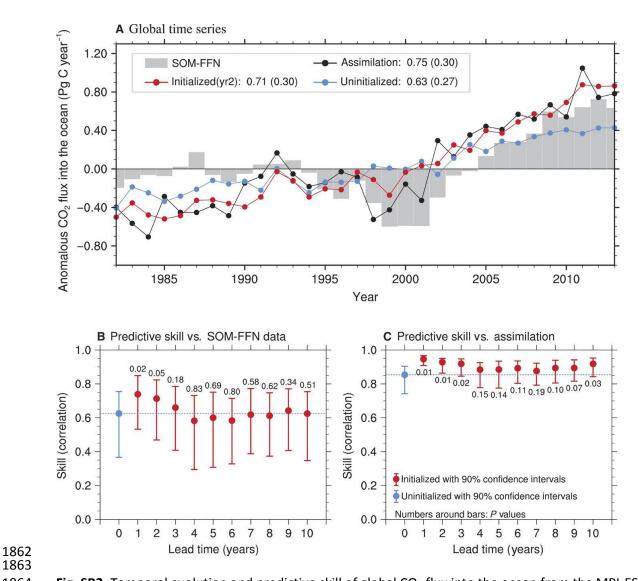
1838 Fig. 9. Global averages of correlations between hindcast and observed precipitation anomalies over the 1839 80°S to 80°N latitudinal band for weeks 1-4 for S2S project models initialized from November to March, 1840 1999–2009. Left: Hindcast quality assessment based on ensemble means using the full ensemble size for 1841 each model, as indicated in the figure legend. Right: Hindcast quality assessment based on ensemble 1842 means using three ensemble members for each model. The reduced spread of the lines shown in the 1843 right panel when ensemble sizes are identical compared to the spread of the lines shown in the left 1844 panel demonstrates the value of the use of larger ensembles for subseasonal precipitation forecasting. 1845 (Adapted from de Andrade et al. 2019.)



- 1846
- 1847 **Fig. 10.** Schematic illustration of relationships between a S2S forecast range of 10-30 days and other
- 1848 prediction timescales, including examples of actionable information that can enable decision making by
- 1849 various sectors. Indicated actions are examples that are not exclusive to a particular forecast range.
- 1850 (After White et al. 2017.)







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1864 Fig. SB2. Temporal evolution and predictive skill of global CO₂ flux into the ocean from the MPI-ESM-HR 1865 decadal prediction system. (A) Annual values of anomalous CO₂ flux into the ocean from data-based 1866 estimates (SOM-FFN; gray) and MPI-ESM uninitialized simulations (blue), year 2 of initialized decadal 1867 predictions (red) and data-constrained assimilation run (black). Anomaly correlations and root-mean-1868 square errors (in parentheses) verifying against SOM-FFN data are indicated. (B) Anomaly correlation 1869 skill for global CO₂ flux into the ocean, verifying against SOM-FFN. The blue dot and dashed line show 1870 the uninitialized skill for which lead time is not relevant, and the red dots initialized skill for different 1871 forecast years, with 90% confidence intervals and P values based on a bootstrap approach indicated. 1872 (C) Like (B), but verifying against the assimilation run. (After Li et al. 2019.)